

### Summary

In the northern Gulf of Mexico, excess dissolved inorganic nitrogen (N) and phosphorus (P) loads from the Mississippi-Atchafalaya River system promote high primary production and contribute to the seasonal development of hypoxic bottom waters on the Louisiana Shelf. While phytoplankton growth is considered to be typically N-limited in marine waters, P limitation has been observed in this region during peak river discharge in spring and early summer. River-induced P limitation is a common phenomenon in coastal hypoxic systems. Although a key aspect of nutrient load reduction strategies for hypoxia mitigation, there is still limited direct evidence for the effect of P limitation on hypoxia. Here we present a synthesis of recent investigations that quantitatively assessed, using a realistic physical-biogeochemical model, the effect of P limitation on primary production and hypoxia development on the Louisiana Shelf.

Simulations show that P limitation delays and displaces westward a portion of river-stimulated primary production and depositional fluxes, resulting in a redistribution of respiration processes toward the western Louisiana Shelf. This redistribution does not promote a westward expansion or relocation of hypoxia. Rather, the onset of hypoxia is delayed and the size of the hypoxic zone reduced. In other words, P limitation dilutes the effects of eutrophication on the Louisiana shelf. Two additive effects explain this reduction, namely the westward shift of organic matter respiration against the backdrop of weakening vertical stratification and the net shift of respiration from the sediments to the water column.

#### Model description **Circulation model** The Regional Ocean Modelling System (ROMS) is configured to simu- The nitrogen cycle model of Fennel et al. (2011) is late the circulation over the Louisia continental shelf. no Cons Co production Oxygen Nitrification Respiration SDet -> LDet PHOSPHORUS INERALIZATIO Sediment oxygen consumption Particulate Organ Particulate Organic Matter

Figure 1. Schematic of the ecosystem model. Black arrows indicate the original nitro gen cycle model of Fennel et al. (2011) and blue arrows the added phosphorus processes (Laurent et al. 2012). Sedimenting particulate organic matter (dashed arrows is instantly remineralized into ammonia and phosphate at the sediment-water interface. Part of the nitrogen is lost into nitrogen gas (N2) through denitrification. Oxygen sources and sinks are represented on the left-hand side panel.

modified to include a dissolved inorganic phosphorus (DIP) compartment (Fig. 2). Phosphorus interacts with the nitrogen cycle model through phytoplankton growth ( $\mu$ ; d<sup>-1</sup>), which is limited by light and temperature ( $\mu(E,T)$ ; d<sup>-1</sup>) and by either dissolved inorganic nitrogen (DIN) or DIP as follows:

$$L_{\text{DIN}} = \frac{NO3}{k_{NO3} + NO3}$$
$$L_{\text{DIP}} = \frac{PO4}{k_{PO4} + PO4}$$
$$\mu = \mu(E, T) \cdot \mathbf{m}$$

Simulations: The model is run for the period 2001 2007. Six simulations are carried out:

- to determine the effect of P limitation.
- concentration in the Mississippi and Atchafalaya rivers.





The simulated annual cycle of nutrient limitation is as follows: no nutrient limitation in winter (Fig. 2a), P limitation in the mid-salinity waters of the Mis- Mar sissippi River plume in early summer (Fig. 2b) and N limitation over the shelf in late summer and fall (Fig. 2c). Open ocean waters are always N-limited. These patterns agree well with the observations (Fig. 2) and bioassays of Sylvan et al. (2006). Interestingly, P limitation does not occur in the Atchafalaya River plume. The shallow depth in this region (<20 m) results in rapid P turnover and N loss through denitrification, which limits the potential for P limitation.

## 2. Shift in primary production

P limitation reduces primary production by 26% near the Mississippi delta between May and July (Fig. 3a). The excess DIN is transported downstream (i.e., westward, Fig. 4), where it fuels primary production in otherwise Nlimited waters (Fig. 3a). This induces a time delay and a westward relocation of part of primary production. Since this relocation occurs over a large area, the overall primary production is diluted over the shelf (see conceptual framework). POM deposition is a quadratic function of primary production in the model and therefore the dilution in primary production results in an overall reduction of POM deposition flux.



# Phosphorus limitation reduces hypoxia in the northern Gulf of Mexico: results from a physical-biogeochemical model

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### **Biological model**

 $\frac{1}{1+NH4/k_{NH4}} + \frac{NH4}{k_{NH4}+NH4}$ 

### $\min(L_{\text{DIN}}, L_{\text{DIP}})$

• The control simulation; results are compared with observations using the L<sub>DIN</sub> and L<sub>DIP</sub> formulations. • A simulation with the original N cycle model

• Additional simulations with ±50% DIP and/or DIN

## **Conceptual framework**

The effect of P limitation on phytoplankton is similar among river-induced eutrophicated systems; P limitation displaces phytoplankton biomass toward downstream waters due to the transport of excess N (Figure 7). However, there is currently no consensus on whether this displacement amplifies or weakens hypoxia. P limitation is generally viewed as a mechanism that relocates or spreads hypoxia and thereby considered detrimental. Any relocation of primary production due to P limitation will also relocate O2 sinks, which could potentially lead to the development of hypoxic conditions in waters that would be normoxic without P limitation.

In flow-through systems, which are characterized by strong freshwater-induced stratification and transport akin to a simple translation along their upstream-downstream axis, a shift of primary production along this axis may well result in a linear effect on hypoxia.

However, excess nutrients in river plumes are being diluted when plumes interact with coastal circulation forced by topography, winds and tides. In these open systems, a "downstream" relocation may spread elevated primary production over a larger area while lowering the maxima of primary production in the affected area, in effect "diluting" the imprint of eutrophication (Fig. 7). Hypoxia is reduced upstream, but does not increase significantly downstream, resulting in an overall reduction of the hypoxic area.



Figure 7. Conceptual model showing the spatial effects along an upstream-downstream transect of river-induced P limitation on DIN concentration (red), phytoplankton biomass (green) and hypoxia (black horizontal lines, top) in an open systems such as the Louisiana Shelf. The difference in phytoplankton biomass in flow-through systems is also indicated (grey line). The type of resource limitation is represented at the top. Along the transect, N is partly removed by denitrification in the sediment. As an indication, spatial distributions are also represented for systems where P is not limiting (dashed lines).

Anomaly (mmol  $O_2 m^{-2} d^{-1}$ )

<u>-20 -15 -10 -5 0 5 10 15 2</u>

#### Primary production

B



Figure 3. Anomaly plots showing spatial and temporal changes in primary production (A), water column respiration (B) and SOC (C) due to P limitation on the Louisiana shelf. Values represent monthly and latitudinal averages for the Louisiana shelf ( $z \le 50$  m). Anomalies are calculated as the results of the simulation with DIP minus the results from the simulation with the N cycle model only.



Journal of Geophysical Research: Oceans, 118(2), 990–1002. Hetland and DiMarco (2008). How does the character of oxygen demand control the structure of hypoxia on the Texas–Louisiana continental shelf? Journal of Marine Systems, 70, 49–62. Laurent et al (2012). Simulating the effects of phosphorus limitation in the Mississippi and Atchafalaya River plumes. Biogeosciences, 9(11), 4707–4723.



Fig 1.

#### Sediment O<sub>2</sub> consumption





### 3. The dilution effect

Changes in bottom water O2 associated with P limitation are asymmetrical over the Louisiana Shelf with a significant increase on the eastern shelf, but only a small decrease on the western shelf (Figure 5). This spatial asymmetry is explained by two additive effects, namely the westward shift of organic matter respiration against the backdrop of weakening vertical stratification and the net shift of respiration from the sediments to the water column.

The intensity of water column stratification varies along the freshwater gradient of the Mississippi River plume. Water column stratification is strongest on the eastern shelf and decreases toward the western shelf, away from the Mississippi River delta (Hetland and DiMarco 2008). Since bottom O2 concentration is highly correlated with stratification intensity on the Louisiana shelf (Fennel et al. 2013), a westward shift in organic matter respiration results in a reduction of hypoxia. In addition, P limitation leads to a redistribution of respiration between sediment and water column (Figure 3b,c). The westward shift of primary production occurs on a broadening shelf thus spreading primary production over a larger area and essentially diluting phytoplankton and detritus. Smaller concentrations of phytoplankton and suspended detritus reduce coagulation and POM deposition; hence more organic matter is respired in the water column and less in the sediment. This leads to an asymmetric effect of P limitation on respiration between the eastern and the western Louisiana shelf (Figure 3b,c). Given the role of SOC on hypoxia development on the Louisiana Shelf (Fennel et al. 2013), the net shift of respiration from sediments to the water column reduces the overall extent of hypoxia (by 29% on average; Laurent and Fennel 2014).

### 4. N and P load mitigation

Four nutrient reduction scenarios were tested to assess their effect on summer hypoxia (Laurent & Fennel 2014): 50% decrease of river DIN (-N), 50% decrease of river DIP (-P), 50% decrease of river DIN and DIP (-NP) and a 50% increase in river DIN with a simultaneous 50% reduction of river DIP (+N-P). The dual N and P load reduction (-NP) maximizes the decrease in hypoxia size and duration (Figure 6). Decreasing N load only (-N) reduces hypoxia on the western shelf but the hypoxic area is larger than with a dual nutrient reduction. Reducing P only (–P) reduces the hypoxic area but not hypoxia duration on the western shelf. Finally, decreasing P but with a simultaneous increase of N (stronger P limitation) lead to a small reduction of the hypoxic area, but also to a longer hypoxia duration on the western shelf (Figure 6). A dual N and P load reduction strategy is therefore recommended to mitigate hypoxia on the Louisiana Shelf.

Fennel et al (2011). A coupled physical-biological model of the Northern Gulf of Mexico shelf: model description, validation and analysis of phytoplankton variability. Biogeosciences, 8(7), 1881–1899. Fennel et al (2013). Sensitivity of hypoxia predictions for the northern Gulf of Mexico to sediment oxygen consumption and model nesting.

Laurent and Fennel (2014). Simulated reduction of hypoxia in the northern Gulf of Mexico due to phosphorus limitation. Elementa: Science of the Anthropocene, 2(1), 000022. Sylvan et al (2006). Phosphorus limits phytoplankton growth on the Louisiana shelf during the period of hypoxia formation. Environmental Science & Technology, 40(24), 7548–53.



showing the relative change in summer 2004 hypoxic area and hypoxia duration at an eastern and western shelf station, associated with a modification in nutrient river load as follow: -50% DIN (-N), -50% DIP (-P), -50% nutrients (-NP) and +50% DIN, -50% DIP (+N-P).



